

# Lecture 20. Temperature Lidar (4)

- ❑ Review of Rayleigh Integration Technique
- ❑ High-Spectral-Resolution Lidar & Rayleigh Doppler Technique
- ❑ Rotational Raman Technique
- ❑ DIAL Temperature Profiling
- ❑ Comparison of Temperature Lidar Techniques
- ❑ Newly Developing Temperature Lidars  
(K Doppler, Solid-state Na and Fe Doppler)
- ❑ Summary

# Review of Integration Technique

□ **Rayleigh integration technique** is applicable in the regions free of aerosols, clouds, and resonance fluorescence, i.e., where the lidar returns are dominated by pure Rayleigh scattering signals. When using non-resonant wavelength, e.g., 532 or 355 nm, it is about 30–100 km, depending on the product of power, aperture, cross-section and efficiency.

$$dP(z) = -\rho(z)g(z)dz + P(z) = \frac{\rho(z)RT(z)}{M(z)}$$



$$T(z) = T(z_0) \frac{n(z_0)}{n(z)} + \frac{M(z)}{R} \int_z^{z_0} \frac{n(z')g(z')}{n(z)} dz'$$

**Number density ratio (relative number density)  $\Rightarrow$  Temperature**

$$\text{var}[T(z)] \approx \frac{T^2(z)}{N_R(z)} + \left\{ \text{var}[T(z_0)] + \frac{T^2(z_0)}{N_R(z_0)} \right\} \exp[-2(z_0 - z)/H]$$

**Large  $N_R(z) \Rightarrow$  smaller temperature errors**

# High-Spectral-Resolution Lidar

□ In the lower atmosphere when aerosols present, the lidar returns contains a narrow spike near the laser frequency caused by aerosol scattering riding on a Doppler broadened molecular scattering profile.

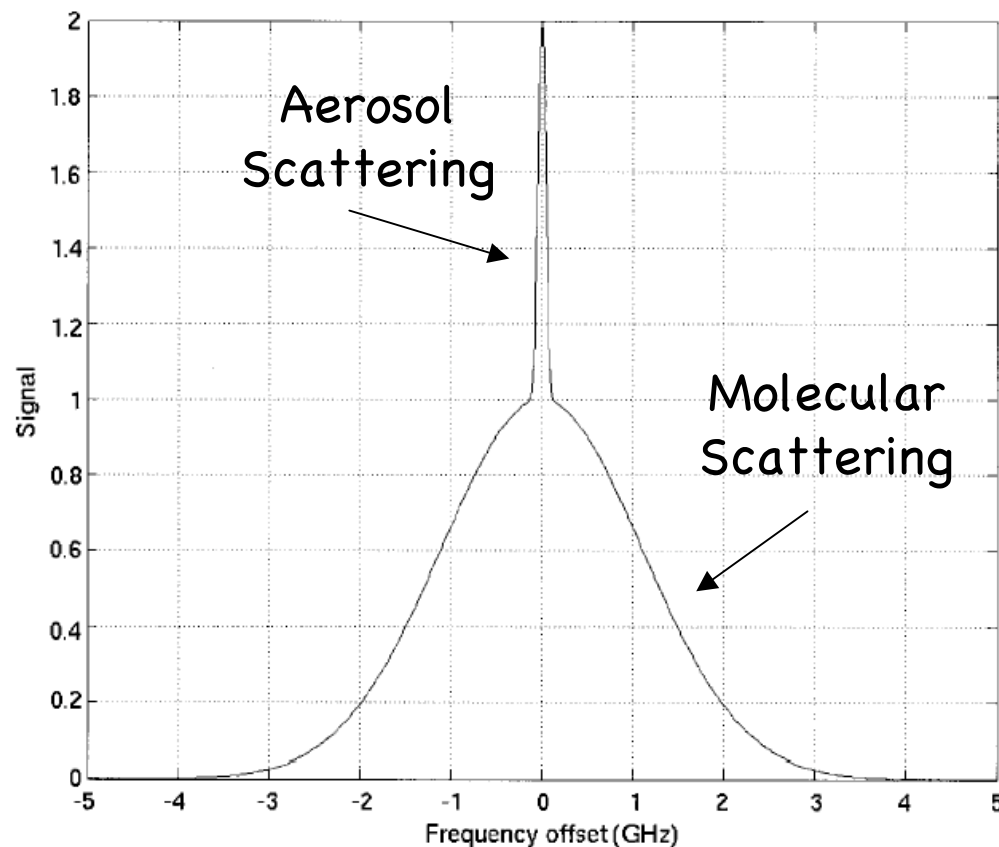


Fig. 5.1. Spectral profile of backscattering from a mixture of molecules and aerosols for a temperature of 300 K. The spectral width of the narrow aerosol return is normally determined by the line width of the transmitting laser.

At  $T = 300$  K, the Doppler broadened FWHM for Rayleigh scattering is 2.58GHz, not 1.29GHz.

Why?

Because Rayleigh backscatter signals have 2 times of Doppler shift!


Courtesy of Dr. Ed Eloranta  
University of Wisconsin

# Doppler Shift in Rayleigh Scattering

- Refer to textbook 5.2.2.4 Lidar wind vs radar wind measurements

Momentum Conservation  $m\vec{v}_1 + \hbar\vec{k}_1 = m\vec{v}_2 + \hbar\vec{k}_2$

Energy Conservation  $\frac{1}{2}mv_1^2 + \hbar\omega_1 = \frac{1}{2}mv_2^2 + \hbar\omega_2$




$$\omega_1 = \omega_2 + \vec{k}_1 \cdot \vec{v}_1 - \vec{k}_2 \cdot \vec{v}_2 + \frac{\hbar k_1^2}{2m} - \frac{\hbar k_2^2}{2m}$$

- For Rayleigh or radar backscatter signals, we have

$$\vec{k}_2 \approx -\vec{k}_1 \quad \vec{v}_2 \approx \vec{v}_1$$

- The frequency shift for Rayleigh or radar backscattering is

$$\Delta\omega_{\text{Rayleigh, backscatter}} = \omega_2 - \omega_1 = -2\vec{k}_1 \cdot \vec{v}_1$$


# Doppler Broadening in Rayleigh Scatter

- To derive the Doppler broadening, let's write the Doppler shift as

$$\omega = \omega_0 \left( 1 - \frac{2v_R}{c} \right) \quad \longrightarrow \quad v_R = \frac{\omega_0 - \omega}{2\omega_0 / c}$$

- According to the Maxwellian velocity distribution, the relative probability that an atom/molecule in a gas at temperature  $T$  has its velocity component along the line of sight between  $v_R$  and  $v_R + dv_R$  is

$$P(v_R \rightarrow v_R + dv_R) \propto \exp(-Mv_R^2 / 2k_B T) dv_R$$

- Substitute the  $v_R$  equation into the Maxwellian distribution,

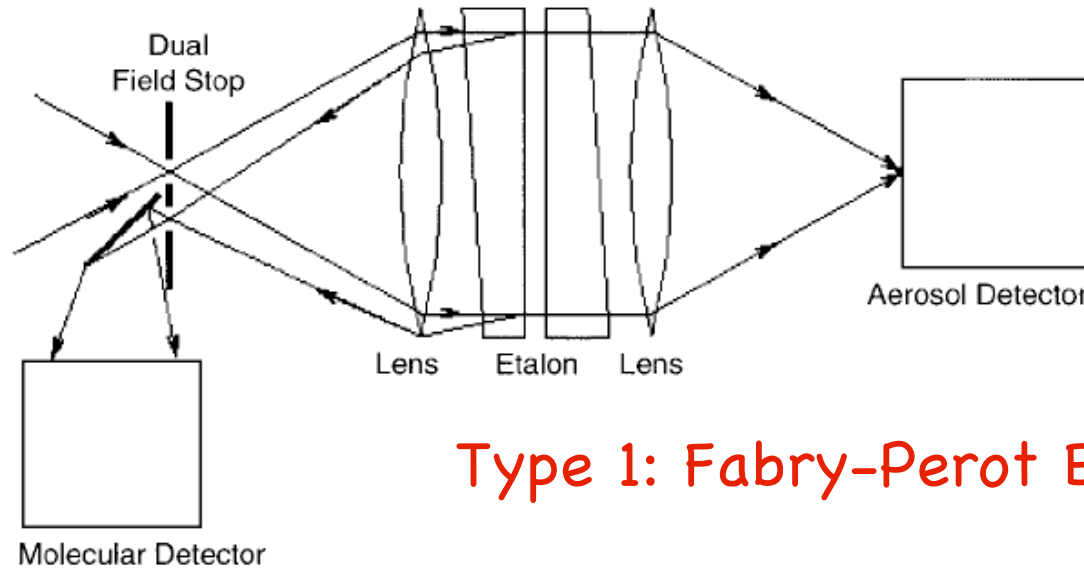
$$I \propto \exp\left(-\frac{M(\omega_0 - \omega)^2}{2k_B T (2\omega_0 / c)^2}\right) (c / 2\omega_0) d\omega$$

- Therefore, the rms width of the Doppler broadening is

$$\sigma_{rms} = 2\omega_0 / c \sqrt{k_B T / M} = \frac{2}{\lambda_0} \sqrt{k_B T / M}$$

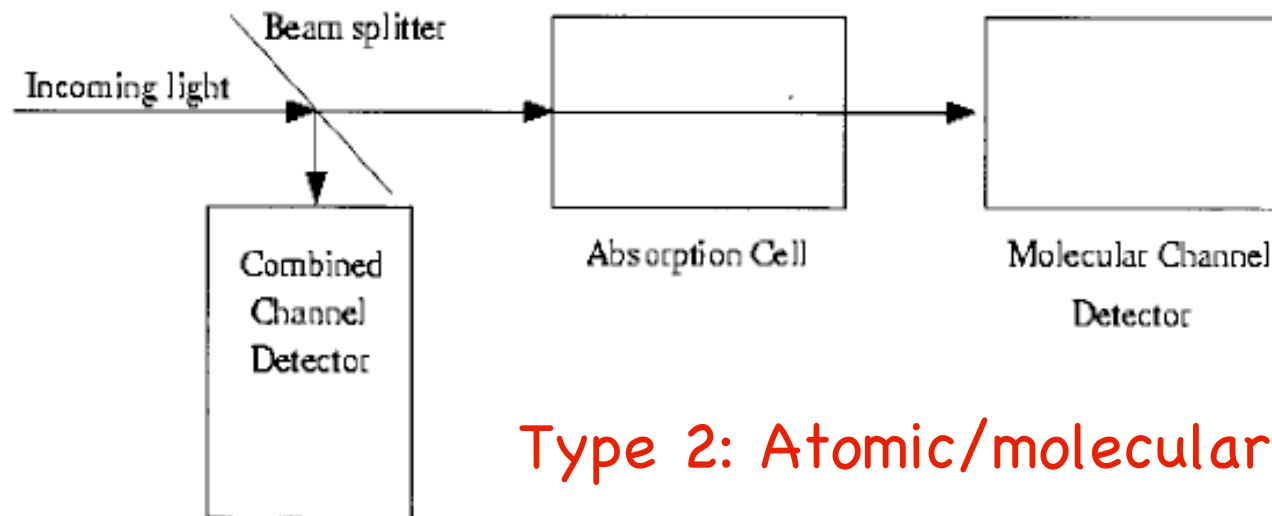
2 times !

# HSRL Receiver Filters



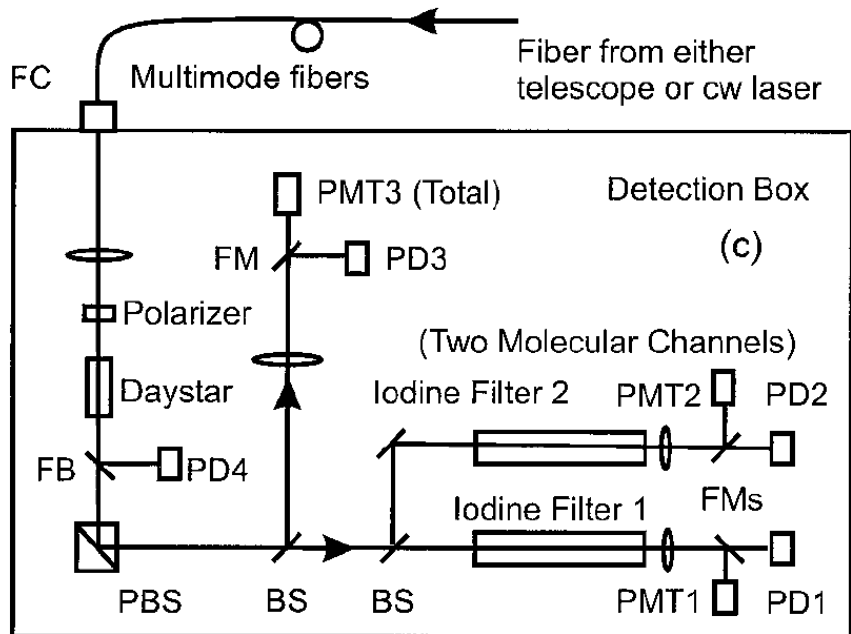
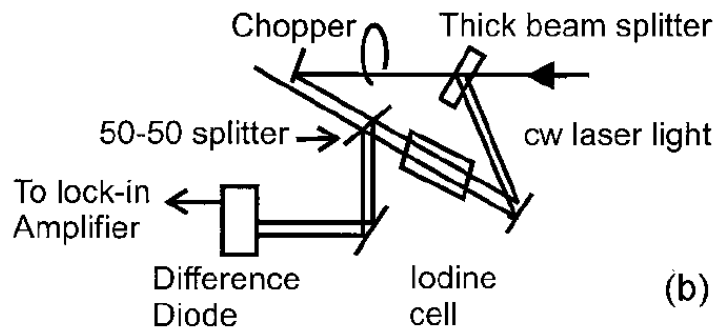
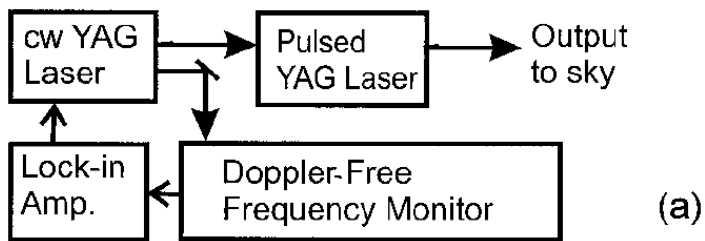
1. Aerosol scattering pass the transmission band
2. Molecular scattering is reflected outside the transmission band

## Type 1: Fabry-Perot Etalon/Interferometer

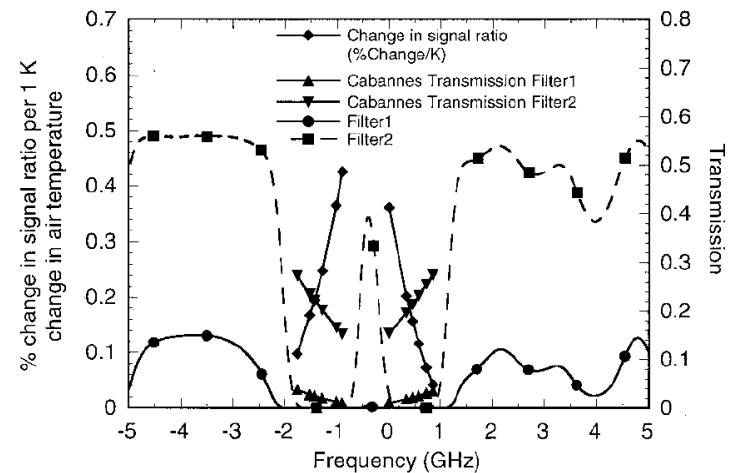


Atomic or molecular absorption block the aerosol scattering

## Type 2: Atomic/molecular blocking filter



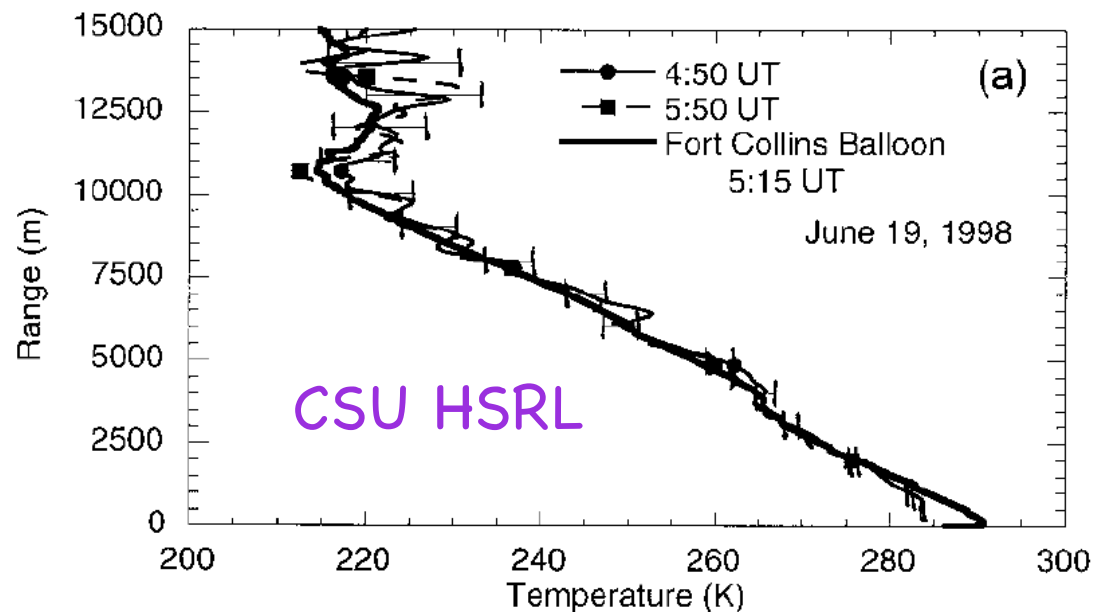
# HSRL Instrumentation Using I<sub>2</sub> Filter



CSU HSRL  
Hair, She, et al. [2001]

# HSRL Temperature Measurements

- ❑ The ratio of the Rayleigh scattering signals passing through two vapor cell filters operating at different temps is a function of atmosphere temperature.
- ❑ Laser has to be single frequency and locked to the narrowband filter. Measurements can go to 15 km.

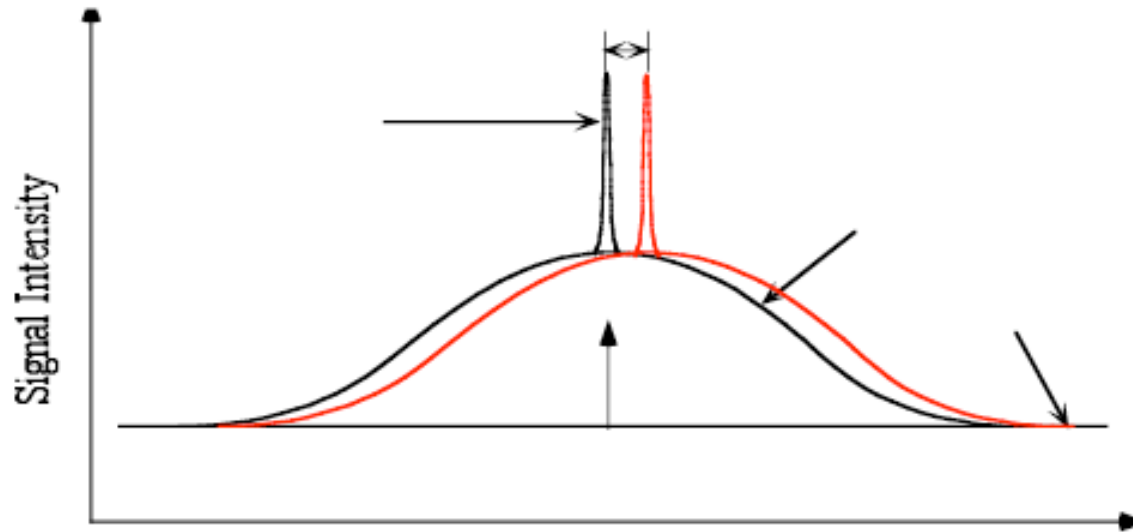


- ❑ Majority of the Rayleigh scattering is filtered out!



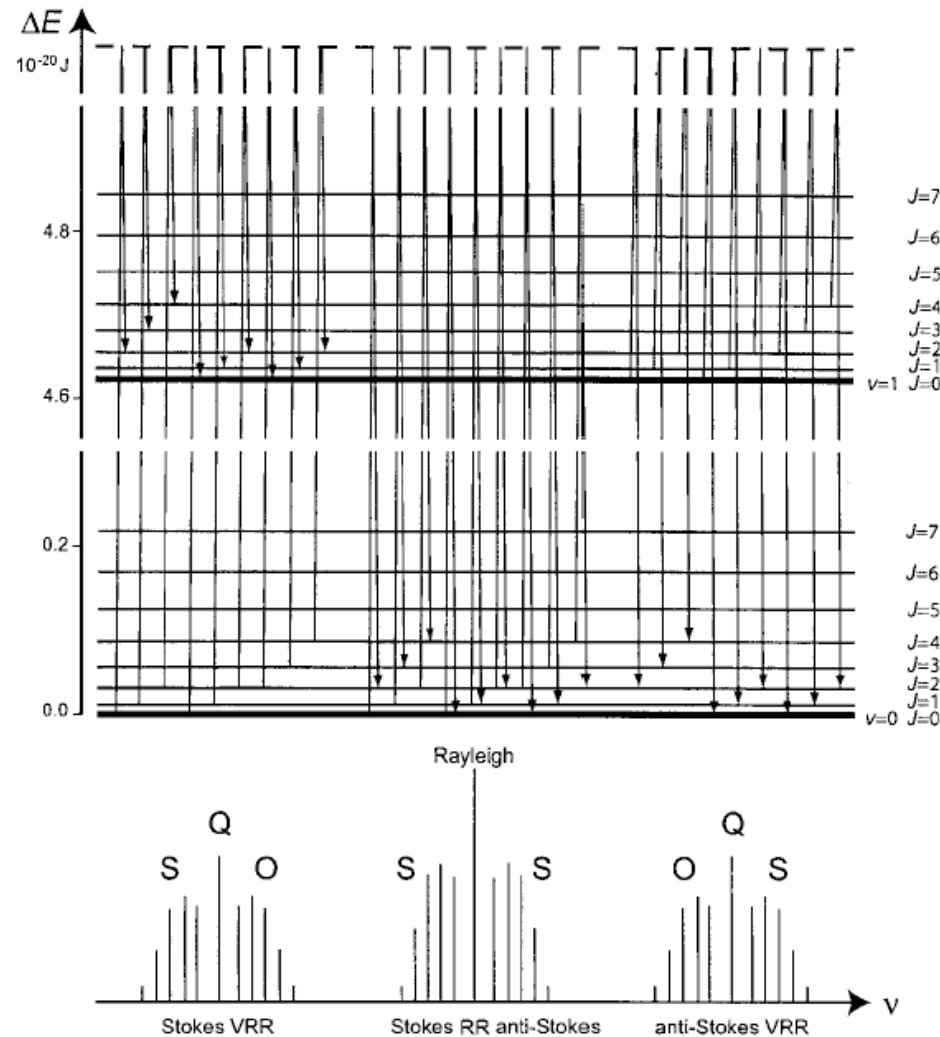
# Rayleigh Doppler Lidar

- ❑ Rayleigh Doppler lidar uses the same Doppler effect of molecular scattering - again, 2 times of the Doppler shift and Doppler broadening!
- ❑ One approach is to use high resolution F-P etalon to image the lidar returns, i.e., turn spectral distribution to spatial distribution.



- ❑ Current issues: suffer low signal levels above 50 km because of decreasing atmospheric density

# Raman Scattering of N<sub>2</sub> and O<sub>2</sub>



**Fig. 9.1.** Vibration-rotation energy levels of the N<sub>2</sub> molecule, Raman transitions, and resulting spectrum.

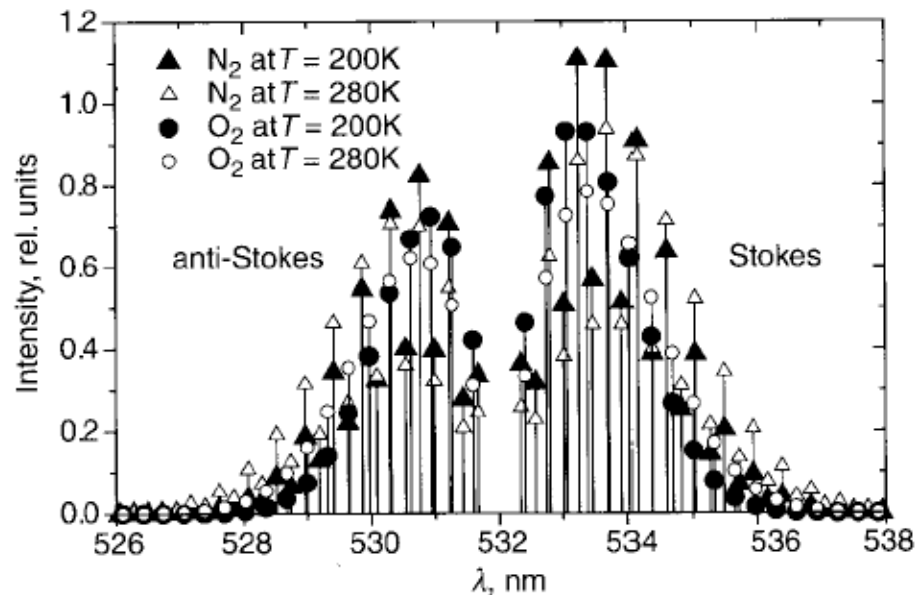
□ Raman shift amount is independent of incident laser wavelength

# Raman Scattering

- Volume backscatter coefficient for single Raman lines

$$\left(\frac{d\sigma}{d\Omega}\right)_J^{\text{RR, VRR}} = k_{\tilde{\nu}}(\tilde{\nu}_1 \mp |\Delta\tilde{\nu}|)^4 \frac{g_N \Phi_J}{Q} \exp\left[-\frac{B_i h c_0 J(J+1)}{k_B T}\right]$$

Which is the product of the transition probability and the population on the initial energy state. So the temperature dependence comes from the population distribution - Boltzmann distribution law!

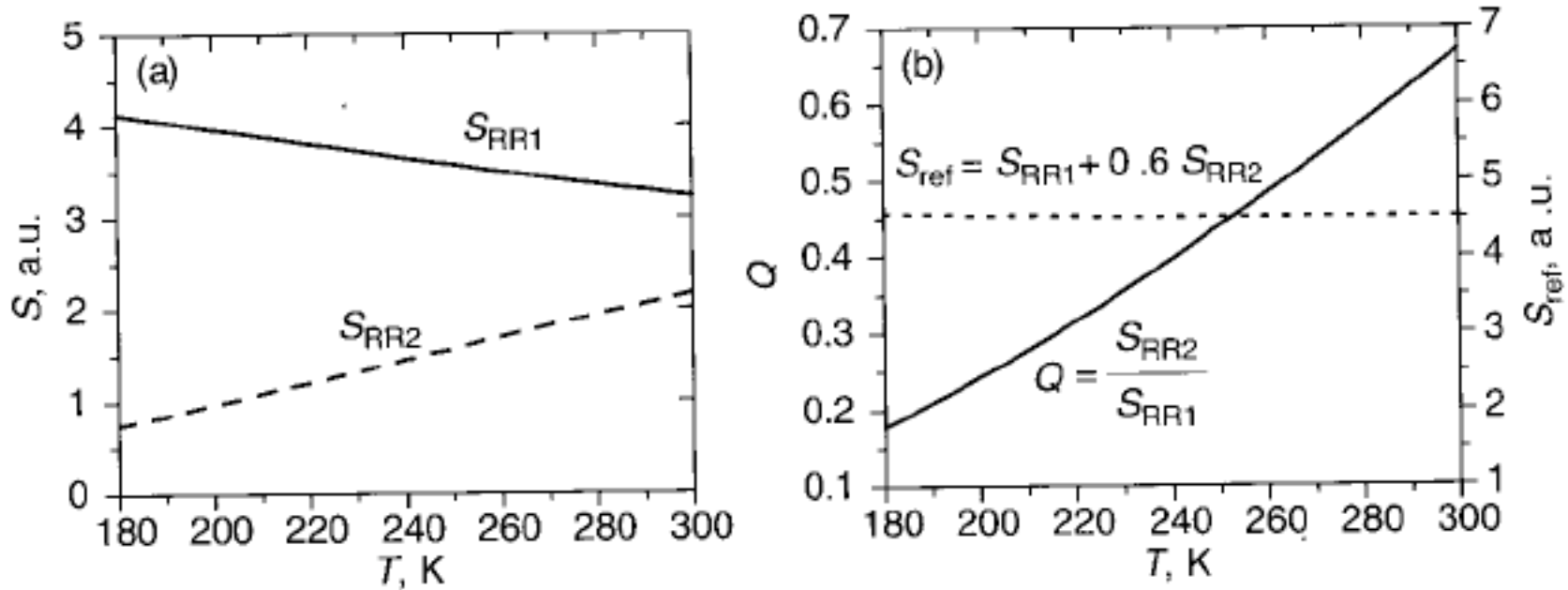


# Rotation Raman Lidar

- ❑ Depending on what the initial energy state is, the line intensity can increase or decrease when temperature increases.
- ❑ If the initial energy state is one of the upper levels of the ground state, increasing in temperature will increase the population on the initial state, so the Raman line intensity will increase.
- ❑ If the initial energy state is the lowest level of the ground state, increasing temperature will decrease the population on the initial state, so the Raman line intensity will decrease.
- ❑ By measuring the intensity of two Raman lines with opposite temperature dependence, the ratio of these two lines is a sensitive function of atmospheric temperature.

$$Q(T, z) = \frac{S_{RR2}(T, z)}{S_{RR1}(T, z)}$$

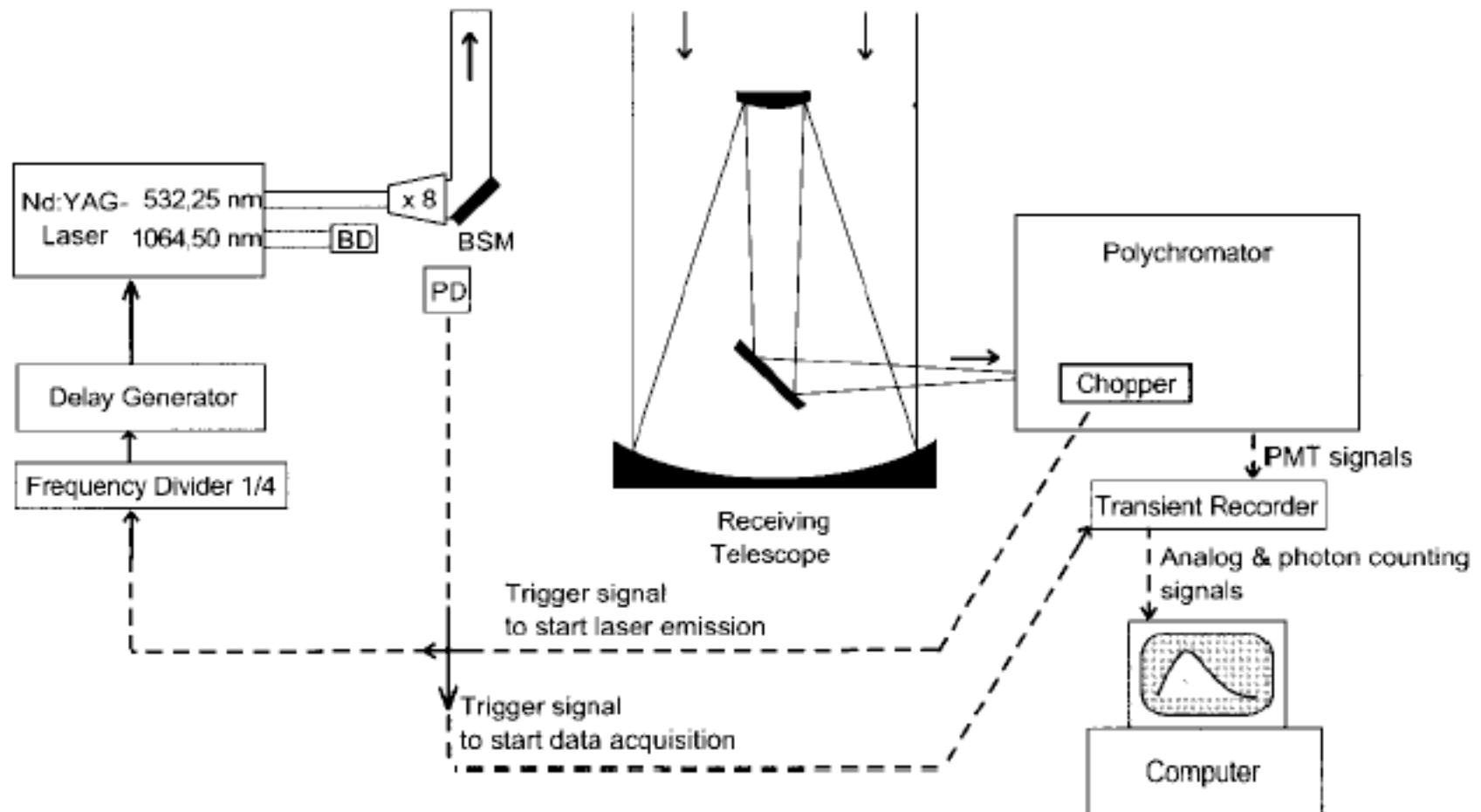
# Rotation Raman Lidar



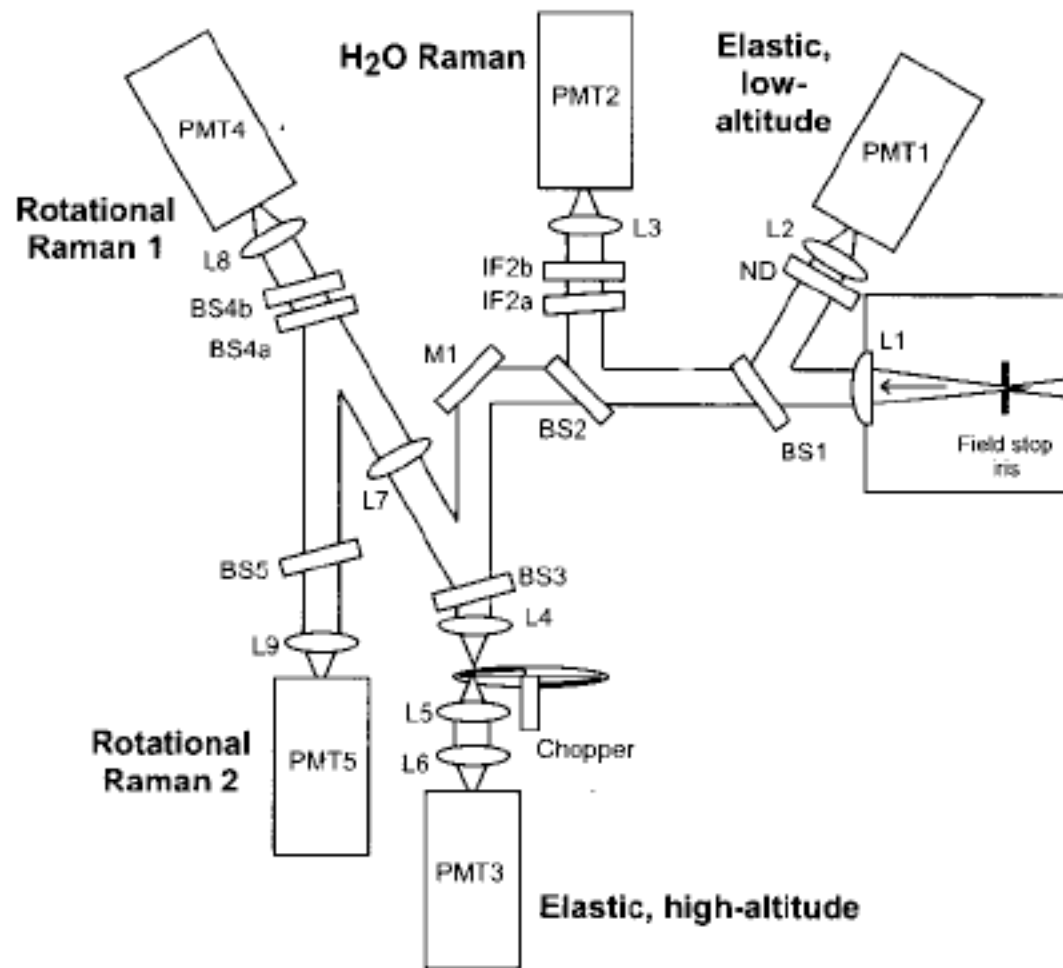
$$Q(T) = \frac{\sum_{i=O_2, N_2} \sum_{J_i} \tau_{RR2}(J_i) \eta_i \left( \frac{d\sigma}{d\Omega} \right)_{\pi}^{RR,i}(J_i)}{\sum_{i=O_2, N_2} \sum_{J_i} \tau_{RR1}(J_i) \eta_i \left( \frac{d\sigma}{d\Omega} \right)_{\pi}^{RR,i}(J_i)}$$

□ Therefore, temperature can be derived from the ratio of two pure Rotational Raman line intensity. This is essentially the same principle as Boltzmann temperature technique!

# Combined Rotation Raman and Elastic Scattering Lidar

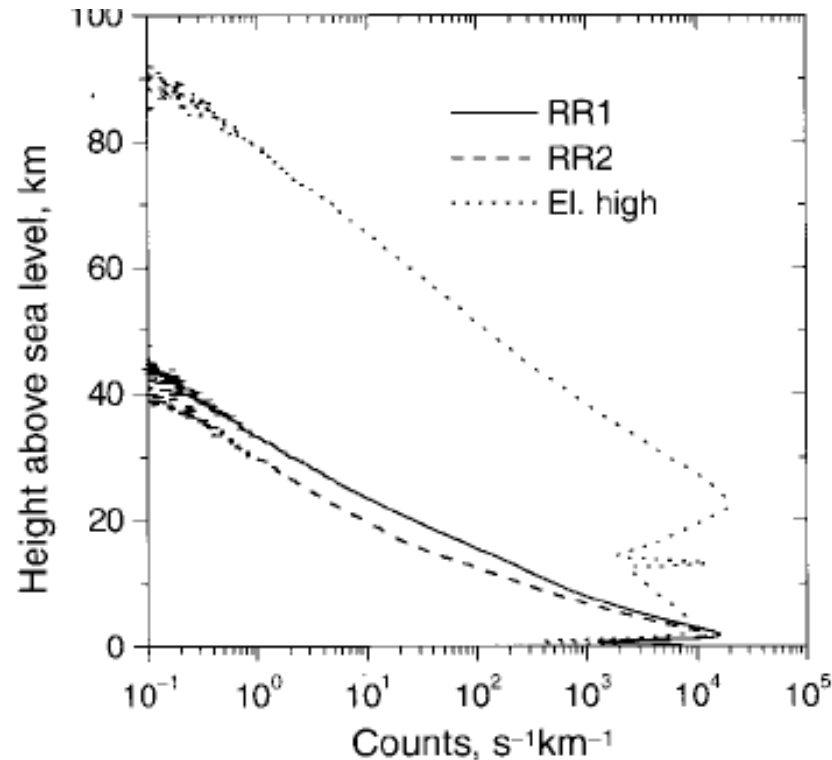


# Rotation Raman + Elastic Lidar



Lidar Polychromator

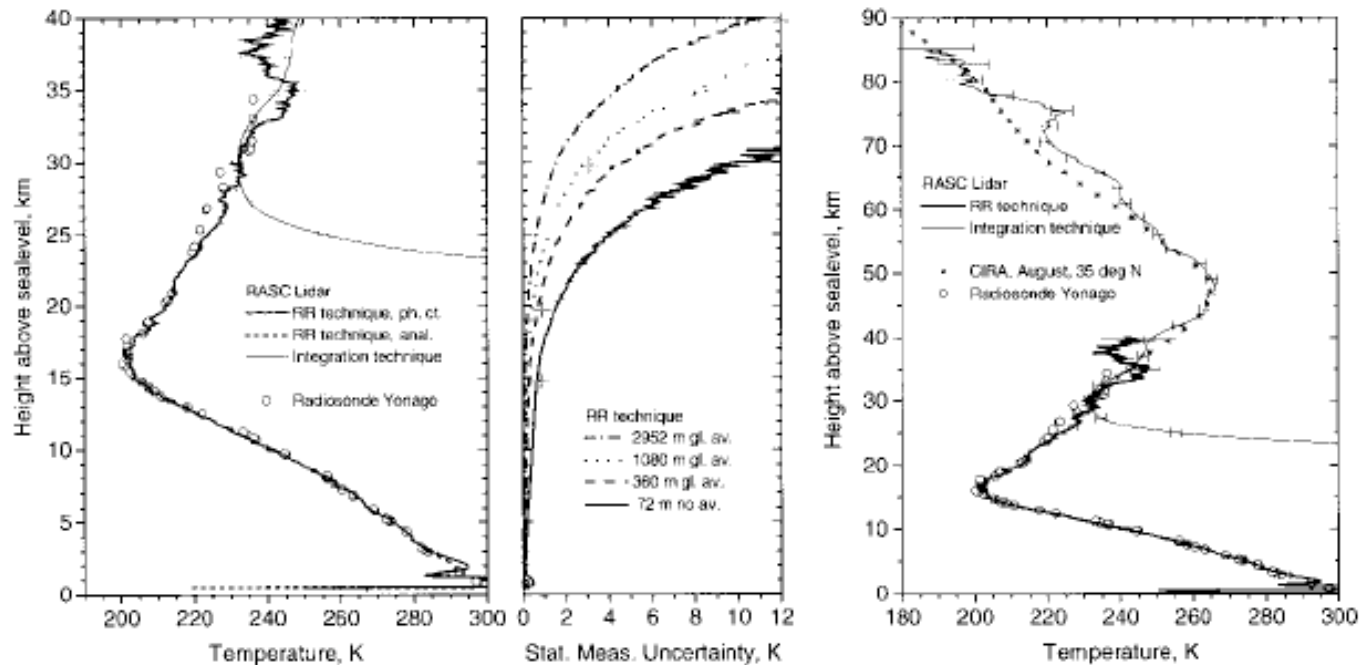
# Results from Combined RR and Elastic Scattering Lidar



**Fig. 10.11.** Intensities of the RASC lidar signals for the temperature measurements: rotational Raman signals (RR1 and RR2) and high-altitude elastic signal (El. high). For this plot, 72 minutes (216,000 laser pulses) of nighttime lidar data were taken with a height resolution of 72 m, summed, the background was subtracted, and the data were finally smoothed with a sliding average of 360 m. The photon emission rate of the laser is  $\sim 8 \times 10^{19}$  photons/s. In the high-altitude elastic signal, the effect of the chopper can be seen below  $\sim 25$  km and the signature of a cirrus cloud in  $\sim 13$  km height.



# Results from Combined RR and Elastic Scattering Lidar



**Fig. 10.12.** Simultaneous temperature measurements with rotational Raman technique and with integration technique (signals see Fig. 10.11). Profiles of a climatological model atmosphere (CIRA-86 for 35°N and the month of the lidar measurements) and of a radiosonde are shown for comparison. Rotational Raman temperature data: height resolution of 72 m up to 15 km height, 360 m between 15 and 20 km height, 1080 m between 20 and 30 km height, and 2952 m above 30 km. Height resolution of the integration technique data is 2952 m. Error bars show the 1- $\sigma$  statistical uncertainty of the measurements [48].

# DIAL Temperature Profiling

- ❑ Molecular absorption coefficient is temperature dependent: both the line strength and the lineshape are function of temperature.
- ❑ So by measuring the molecular absorption coefficient, it is possible to derive temperature if the molecular number density is known. For this purpose,  $O_2$  is chosen because of its constant mixing ratio up to high altitude and suitable absorption lines.
- ❑ In the choice of suitable absorption line, a trade-off must be made between the high temperature sensitivity of the absorption cross-section (high for high initial energy state) and the suitable magnitude of absorption coefficient.
- ❑ Absorption coefficient is also dependent on pressure, making the temperature derivation more difficult.

# Comparison of Temperature Technique

Technique	Lidars	Applications
<b>Doppler Technique:</b> <b>temperature dependence of Doppler broadening</b> (1 time Doppler shift and Doppler broadening for single absorption or emission process) (2 times Doppler shift and Doppler broadening for Rayleigh scattering)	Resonance fluorescence Doppler Lidar: Doppler broadening and Doppler shift of resonance fluorescence absorption cross-section (scan and ratio techs)	Mesosphere and Lower Thermosphere temperature and wind (75-120 km)
	Rayleigh Doppler Lidar : Doppler broadening of molecular scattering	Lower mesosphere, stratosphere and troposphere temperature and wind (up to 60 km)
	High-Spectral-Resolution Lidar: Doppler broadening of molecular scattering, ratio of two signals	Stratosphere and troposphere temperature and wind (up to 30 km)
<b>Boltzmann Technique:</b> <b>temperature dependence of population ratio</b>	Resonance fluorescence Boltzmann Temperature Lidar: population ratio on the lowest two ground states	Mesosphere and Lower Thermosphere temperature (75-120 km)
	Rotation Raman Temperature Lidar: ratio of two Raman line intensities and population on different initial energy states	Troposphere and stratosphere temperature
<b>Integration Technique:</b> <b>hydrostatic equilibrium and ideal gas law</b>	Rayleigh Integration Temperature Lidar: atmospheric density ratio to temperature, integration from upper level	Stratosphere and mesosphere temperature (30-90 km)
<b>DIAL</b>	Differential Absorption Lidar: Temp-dependence of line strength and lineshape	Boundary layer temperature

# Comparison of Temperature Technique

- ❑ All are temperature-dependent effects and phenomena !
- ❑ Applications of similar techniques in other remote sensing techniques -
  - ❑ Doppler effect: FPI for wind and temp measurements
  - ❑ Boltzmann distribution / rotational Raman: All-sky-imager, Bomem temp mapper
  - ❑ Integration technique: rocket falling sphere, cw searchlight, satellite drag, etc.

# Solid-State K Doppler Lidar

□ See Lecture 18

# Solid-State Na Doppler Lidar

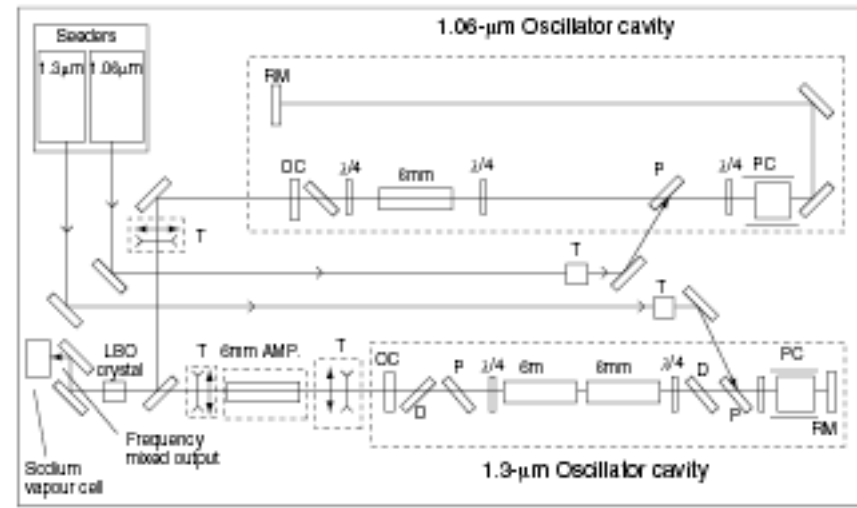
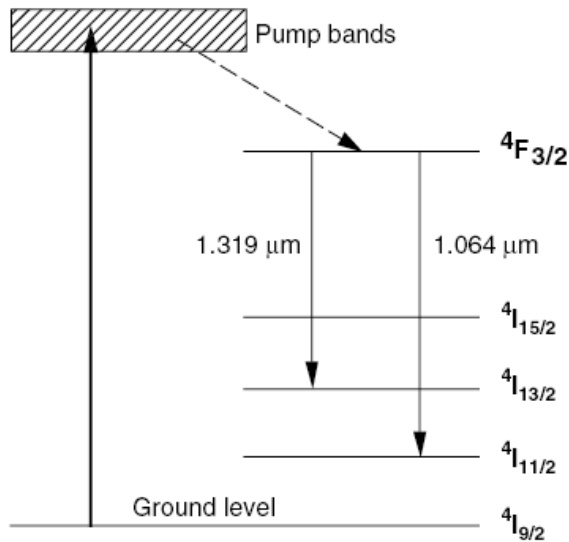
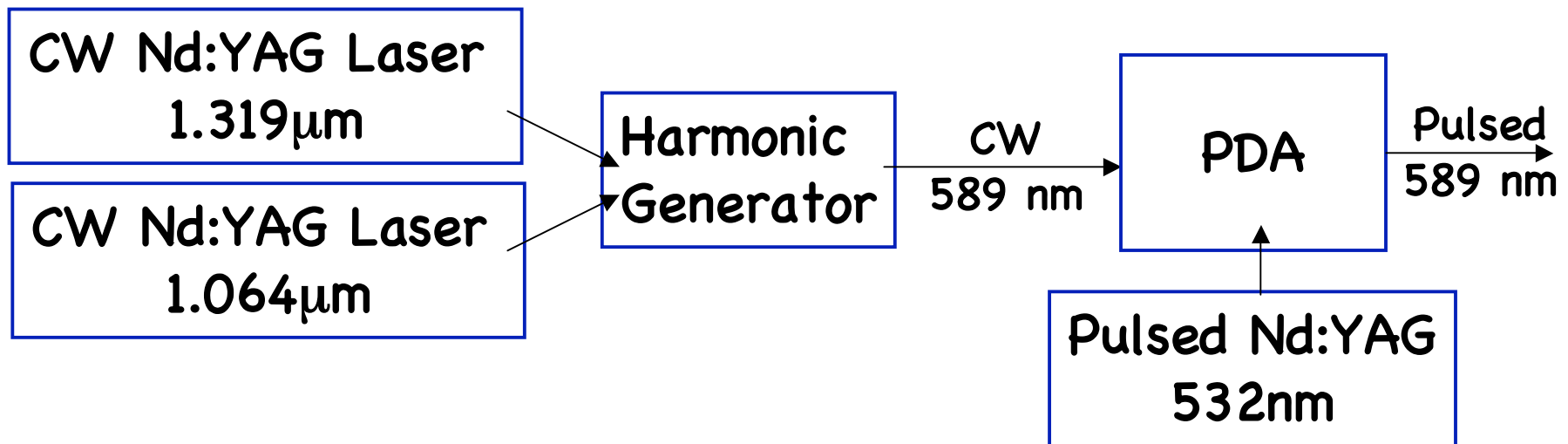
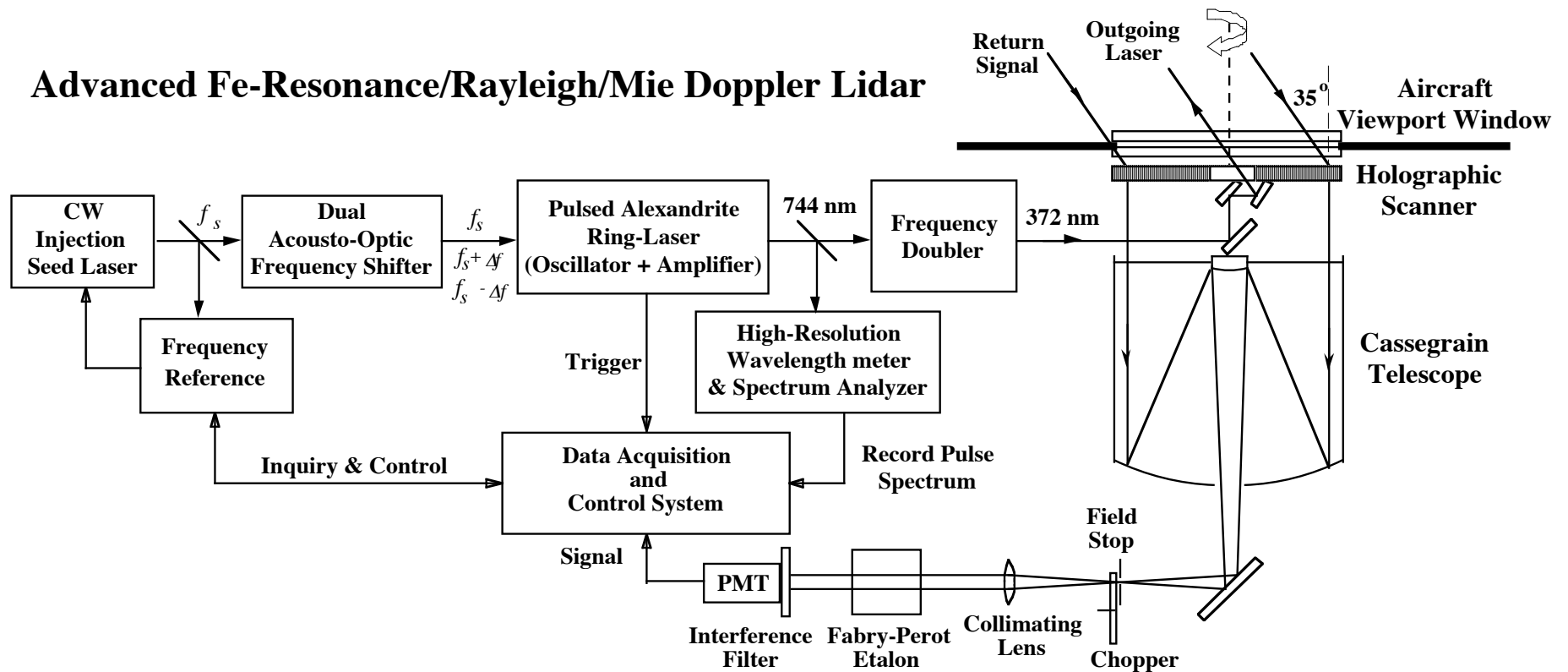


Figure 5.36 Energy level diagram for Nd:YAG laser.



# Proposed Solid-State Fe Doppler Lidar

## Advanced Fe-Resonance/Rayleigh/Mie Doppler Lidar



# Summary

- ❑ Temperature-dependent effects and phenomena
- ❑ Doppler Technique
- ❑ Boltzmann Technique
- ❑ Integration Technique
- ❑ Rotational Raman Technique
- ❑ DIAL Technique
- ❑ High-spectral-resolution Lidar
- ❑ Rayleigh Doppler Lidar